L-EDGE XANES MEASUREMENTS OF THE OXIDATION STATE OF TUNGSTEN IN IRON BEARING AND IRON FREE SILICATE GLASSES. L. R. Danielson^{1,2}, K. Righter², S. Sutton³, M. Newville³, L. Le^{1,2}, ¹Jacobs Sverdrup Co., Houston, TX 77058 United States (lisa.r.danielson@nasa.gov), ²NASA JSC, 2101 NASA Road One, Houston, TX 77058 United States, ³GSECARS University of Chicago, 9700 South Cass Avenue, Bldg. 434A, Argonne, IL 60439 United States.

Introduction: Tungsten is important in constraining core formation of the Earth because this element is a moderately siderophile element (depleted ~ 10 relative to chondrites) and, as a member of the Hf-W isotopic system, it is useful in constraining the timing of core formation. A number of previous experimental studies have been carried out to determine the silicate solubility and metal-silicate partitioning behavior of W, including its concomitant oxidation state. However, results of previous studies (Fig. 1) are inconsistent on whether W occurs as W⁴⁺ or W⁶⁺.

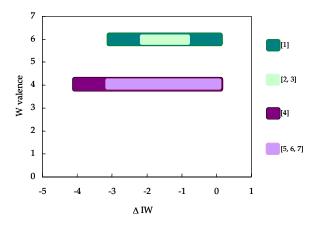


Figure 1. Comparison of W valence from previous results.

It is assumed that W^{4+} is the cation valence relevant to core formation [8]. Given the sensitivity to silicate composition of high valence cations [8], knowledge of the oxidation state of W over a wide range of fO_2 is critical to understanding the oxidation state of the mantle and core formation processes. This study seeks to measure the W valence and change in valence state over the range of fO_2 most relevant to core formation, around IW-2.

Experiments: Two compositions were used to determine the effects of iron content. Initial experiments, reported in [9] were conducted at 1300 °C, for durations of 24 to 96 hours and air or water quenched. One series was conducted using the An-Di eutectic, from QFM to IW-5 (log fO₂ -7.25 to -16. Experiments using an ankaramite starting composition were conducted from IW-1 to IW-5 (log fO₂ -11.75 to -16). Experiments were doped with 1wt% of WO₃. For both starting compositions, at IW-1, one set of experiments was doped with 1wt% of WO₂. Experiments at IW-2 and

above were conducted using a Re wire loop technique, while experiments below IW-2 were conducted in sealed silica tubes as in Fig. 2.

The most recent experiments reported herein expanded the fO_2 range, from in air to IW-10 (ankaramite runs up to log fO_2 -0.4, CMAS runs down to log fO_2 -20) of the data set, and were designed to in an attempt to improve success with analyzable runs at low fO_2 . Run duration was increased to 120 hours and all experiments were water quenched. Because the solubility of W decreases to ~100 ppmw [4] at the low fO_2 conditions of these experiments, amount of W as WO_3 added was below 1000 ppmw.

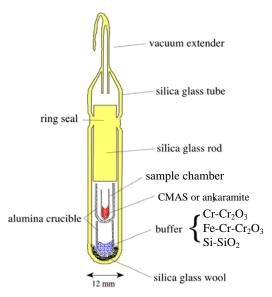


Figure 2. Schematic of sealed silica tube experiments conducted at lowest fO₂, below IW-2.

Analytical: A monochromatic X-ray beam from a Si(111) double crystal monochromator was focused onto the sample and the fluorescent X-ray yield was plotted as a function of incident X-ray energy (more detail can be found in [10]). The oxidation state of tungsten was inferred from the energy of the first peak in the LIII-edge derivative spectrum. WO₂, WO₃, Fe-WO₄, CaWO₄, were used as standards.

Results: Results (Fig. 3 and 4) for both the ironbearing and iron-free starting materials suggest that only W^{6+} is present from the most oxidized conditions to IW (log fO₂ -10.75). At IW, tungsten starts to exhibit mixed valence but is still dominated by W^{6+} for

ankaramite. At IW-2, W⁴⁺ becomes more abundant for ankaramite, but a definitive transition to W⁴⁺ below IW-2 has not been observed. CMAS has a lower W valence than ankaramite below IW, with the most reduced state observed being W⁴⁺. These results suggest that W⁴⁺ would be dominant below IW-2.

Even with improved experimental technique focused on reducing the "nugget effect," at IW-2 and below, metal exsolved from the silicate, complicating the analyses. Analyses in nugget free regions of the silicate portion of the experiments cluster around W^{5+} , while analyses in nugget-rich regions cluster around W^{2+} .

Four CMAS run products from Ertel [4] were also analyzed and shown for comparison in Fig. 4. These samples were reported as W⁴⁺, from the slope of the W concentration vs. log fO₂ line. The Ertel run products are in good agreement with our CMAS run products, showing decreasing W valence with decreasing fO₂, but appear to be mixed W⁶⁺ and W⁴⁺ valence.

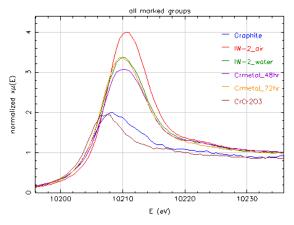


Figure 3. A sampling of results from some CMAS ru product analyses (figure 4), showing range of W valence, from W^{6+} (IW+1 air quench, red line) to around W^{2+} (IW-5, using Cr-Cr₂O₃ buffer, brick line). Data show differences between air (red line) and water quench (green line) at IW-2, with a 0.5+ reduction in W valence observed for a water quench. Other lines indicate a run in a graphite capsule, and runs using only Cr metal as a buffer.

Discussion and Conclusions: Both CMAS and ankaramite glasses show W⁶⁺ above IW and mixed valence below IW. The transition to W⁴⁺ only, appear to happen at or below IW-2 for iron free systems, but probably below IW-3 for iron bearing systems. Analyses of run products from previous experiments [4], support the observation the iron free transition to only W4+ occurs at or below IW-2.

The mixed states below IW-2 may result from analyses in which both silicate glass and exsolved Wbearing metal are present in the analytical volume in varying proportions. This suggests if all W were dissolved in the silicate in a nugget free condition, W valence would be less than W⁵⁺ at lowest fO₂.

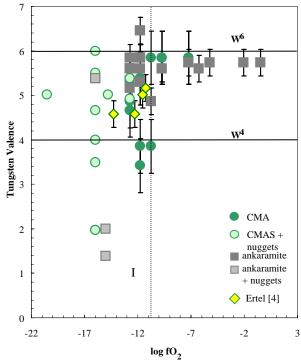


Figure 4. Summary of valence results inferred from the energy of the first peak in the LIII-edge derivative spectrum.

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References: [1] Schmitt et al. (1989) Geochim. Cosmochim. Acta 53(1): 173-185. [2] Walter and Thibault (1995) Science 270: 5239, 1186 – 1189. [3] Hillgren et al. (1996) Geochim. Cosmochim. Acta 60(12), 2257-2263. [4] Ertel et al. (1996) Geochim. Cosmochim. Acta 60(7), 1171-1180. [5] Jones (1998) Meteoritics & Planet. Sci., 33, A79. [6] Lauer and Jones (1999) LPSC XXX, 1617. [7] Wade and Wood (2005), Earth and Planet. Sci. Let., 236(1-2), 78-95. [8] Jaeger and Drake (2000) Geochim. Cosmochim. Acta 64, 3887-3895. [9] Danielson et al. (2006) LPSC XXXVIII, 2113. [10] Sutton et al. (2002) Reviews on Mineralogy & Geochemistry; Appl of Synchrotron Rad in Low-T Geochem & Environ Sci, Fenter, Rivers, Sturchio, Sutton, eds., Min. Soc. Amer., 429 - 483.